



National Centre for Atmospheric Science







Wilcox et al. (2020)





(a) Global surface temperature change relative to 1850–1900

(b) September Arctic sea ice area





www.ncas.ac.uk www.met.reading.ac.uk/~laura/home

AR6WG1 SPM

About me



Associate Professor at the National Centre for Atmospheric Science, University of Reading

Chair of the Regional Aerosol Model Intercomparison Project (RAMIP)

IPCC AR6 WG1 Chapter 8 author (water cycle change)



The Coupled Model Intercomparison Project



Eyring et al. (2016)



Core DECK (piControl,) and historical simulations, with endorsed MIPs exploring key science questions

A MIP explores the same question, in the same (ish) way, in multiple models, allowing us to explore the effects of model **structural uncertainty**

CMIP experiments allow us to explore the effects of emission uncertainty (e.g. ScenarioMIP), or quantify the role of individual climate forcers in the role of climate change (e.g. DAMIP, AerChemMIP), etc.

Considering these experiments as part of a MIP allows us to better quantify uncertainty, and to learn which processes might be important for climate responses to forcing

The CMIP6 models



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78 models from 32 centres participated in CMIP6

Example family tree for 33 CMIP6 models.

Models branching further to the left are more dependent

Labels with the same colour indicate models with obvious dependencies, such as shared components or the same origin

Models with no clear dependencies are labeled in black.

An estimation of internal variability is given using grey shading

Brunner et al. (2020)



The CMIP6 models

👬 OSF**home 🚽**

Opinion: The Role of AerChemMIP in Advancing Climate an...

CMIP6_AerChemMIP_model_details.csv

Sheet 1

Show rows with cells including:

		Emergent			Model com					
Model	Centre	ECS (Schlu	ERF_AA	ASR_HD	Atmosphere	Aerosol	Chemistry	Ocean	Sea Ice	Land
ACCESS-CM2	CSIRO-AR	4.72	-1.09	2.19	MetUM-Had	UKCA-GLO		ACCESS-O	CICE5.1.2	CABL
ACCESS-E	CSIRO	3.87	-1.15	1.90	HadGAM2	CLASSIC (v		ACCESS-O	CICE4.1	CABL
AWI-ESM-1	AWI	3.16			ECHAM6.3	MACv2-SP		FESOM 1.4	FESOM 1.4	JSBA(
BCC-CSM2	BCC	3.04		0.60	BCC_AGC	MACv2-SP		MOM4	SIS2	BCC_
BCC-ESM1	BCC	3.26		2.04	BCC_AGC		BCC-AGCM	MOM4	SIS2	BCC_
CAMS-CSM	CAMS	2.29		-0.01	ECHAM5_C			MOM4	SIS 1.0	CoLM
CanESM5	CCCma	5.62	-0.85	1.61	CanAM5	Interactive	Specified ox	NEMO3.4.1	LIM2	CLAS:
CAS-ESM2-0	CAS	3.51		2.51	IAP AGCM 5.0	IAP AACM	IAP AACM	LICOM2.0	CICE4	CoLM
CESM2-FV2	NCAR	5.14		2.95	CAM6	MAM4	MAM4	POP2	CICE5.1	CLM5
CESM2	NCAR	5.16	-1.37	2.55	CAM6	MAM4	MAM4	POP2	CICE5.1	CLM5
CESM2-WA	NCAR	4.79		2.82	WACCM6	MAM4	MAM4	POP2	CICE5.1	CLM5
CESM2-WA	NCAR	4.75		2.74	WACCM6	MAM4	MAM4	POP2	CICE5.1	CLM5
CIESM (EC	тни	5.67		0.45	CIESM-AM	MAM4	trop_mam4	CIESM-OM	CICE4	CIESN

https://doi.org/10.17605/OSF.IO/8FWJ3



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CIESN	April 10, 20	25					
UESN	April 10, 20	25		Wilco	x (2025)		

User beware....

Different models aren't always so different....

- CESM2 and CESM-FV2 just have different resolutions
- CESM2 and CESM2-WACCM have atmospheric chemistry

The same model is sometimes different....

- GISS use ensemble physics codes to distinguish between model versions.
- rXiXpXfX
- GISS-E2-I-G rlilpIfI, rlilp3fI, and rlilp5fI all have different chemistry schemes very different models, which are often lumped together due to this labelling

While sometimes, the same model really is just the same...

 CanESM5 p1 and p2 refer to a micro perturbation. As long as you're not looking at variables closely related to the perturbed variable, you can just lump all of these together to create a larger ensemble

National Centre for Atmospheric Science Wilcox (2025)

Global aerosol



William Putman, NASA/Goddard

GEOS-5, 10km resolution Red: Dust Blue: Sea salt Green: Smoke

ke White: Sulphate



Aerosol effects on climate



Many ways for aerosols to interact with climate

Some models don't simulate these interactions

Different models have different approaches to representing these interactions

IPCC (2007)



Aerosol effects on climate



Aerosol forcing is the most uncertain anthropogenic climate forcing, reflecting challenges with both observations and modelling

IPCC (2013)



1986 - 2005 mean sulphate load



- Substantial inter-model diversity in aerosol load, even in the preindustrial period
 Diversity in absolute mass and distribution
 - Factor of four spread in the global mean

Wilcox et al. (2015), GRL



2014 vs. 1850 aerosol ERF



 Different aerosol process representation means this diversity propagates through to ERF
 Wilcox et al. (2020), ACP



2014 vs. 1850 GHG ERF



 Different aerosol process representation means this diversity propagates through to ERF



Wilcox et al. (2020), ACP

2014 vs. 1850 anthropogenic ERF



• Different aerosol process representation means this diversity propagates through to ERF Wilcox et al. (2020), ACP



Model diversity can:

- I. show us which processes are important for model performance
- 2. help us to understand the causes of biases in our own model
- 3. help us to understand the physical drivers of uncertainties in the simulated response to forcing
- 4. be used to constrain model estimates

But, it can be difficult to use an 'ensemble of opportunity' to isolate the role of the thing you're interested in



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Model diversity can help us to understand biases in our own model



Earth system models cool too much in the mid-twentieth century

Linked to large sensitivity to aerosol changes in these models

This example is a case of more sophisticated models having poorer performance....

Zhang et al. (2021)



Better is not always better...



Pothole bias not present in physical models

Comparison of ESMs and physical models within the same family shows that the bias relates to excessive high latitude cooling



Zhang et al. (2021)



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AMOC response to aerosol changes

S C SC CE

Decadal variability in the Atlantic Meridional Overturning Circulation (AMOC) is influenced by changes in anthropogenic aerosol, but the **extent and mechanism of influence is uncertain.**

Differences between CMIP5 and CMIP6 attributed to differences in the strength of the aerosol forcing.



Menary et al. (2020)





Not possible to calculate aerosol forcing for all CMIP6 models (needs a dedicated RFMIP experiment) so design a metric that can be calculated from the historical experiment...

ASR_HD: SH - NH net solar radiation at the top of the atmosphere

-> positive values indicate less radiation absorbed by NH



Robson et al. (2022)



ASR_HD: SH - NH net solar radiation at the top of the atmosphere

-> positive values indicate less radiation absorbed by NH

Strong models have a linear change in ASR_HD between 1850 and 1985 **greater than 1.5 Wm**⁻² -> 9 strong models and 8 weak models

Increase in both ASR_HD and the AMOC from 1850– 1985 with the fastest increase over ~1940–1985

Strong models have **4x larger anomaly in ASR_HD**, and **8x larger anomaly in AMOC**, vs. weak models for 1965-1985 relative to 1850-1879.

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SHF: surface heat flux turHF: turbulent heat flux sNetSW: surface net shortwave sNetLW: surface net longwave

SHF anomalies dominated by strong models

SHF dominates the overall surface density flux anomalies

AMOC anomalies in CMIP6 are consistent with the evolution of **surface heat fluxes, and their impact on surface density fluxes, driving the AMOC** in the 'strong' models



SHF: surface heat flux turHF: turbulent heat flux sNetSW: surface net shortwave sNetLW: surface net longwave

MMM SHF dominated by sNetSW

Differences in SHF anomalies between *strong* and *weak* models are **dominated by differences in turHF**

Relationship between AMOC and SHF is **due primarily to turHF**

Robson et al. (2022)





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Constraining simulated AMOC changes



Increased salinity in strong models inconsistent with observations

North American cooling, and weak trends in Northern Hemisphere temperature are inconsistent with observations

Forced AMOC strengthening in strong models not consistent with observations -> aerosol forcing, or the response to it, is too large in strong models

Robson et al. (2022)



Emergent constraints



Identify an empirical relationship between an observable variable and response to forcing, with credible physical mechanisms

Observed range of predictor can then be used to constrain simulated response by e.g. weighting or rejecting models outside observed range

Florent Brient, 2016



Emergent constraints



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Simpson et al. (2021)



Emergent constraints



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Past temperature trend with future temperature trend, and tropical cloud properties with future temperature trend



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HadGEM3-GC3.1 ensemble with scaled aerosol emissions



Scaling	ERF (W/m2)
0.2	-0.3
0.4	-0.6
0.7	-1.0
1.0	-1.3
1.5	-1.6

• Span a large portion of the IPCC range for aerosol ERF





HadGEM3-GC3.1 ensemble with scaled aerosol emissions



 All scalings warm too quickly since 1981 - GC3.1 has a high climate sensitivity and large aerosol forcing
 Dittus et al., 2020



HadGEM3-GC3.1 ensemble with scaled aerosol emissions



The effect of the uncertainty in aerosol radiative forcing on GMA and GMI is a reduction of 2.99 % and 1.93 % respectively, when increasing the scaling across its range

Magnitude of this impact is equivalent to that from a degree of global warming Shonk et al., 2020



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